



# Use of Time Series Reference Sites to Assess Numerical Weather Prediction Reanalyses: Surface Heat Fluxes and Surface Cloud Forcing in the Eastern Tropical Pacific

Meghan Cronin (NOAA Pacific Marine Environmental Laboratory) Meghan.F.Cronin@noaa.gov

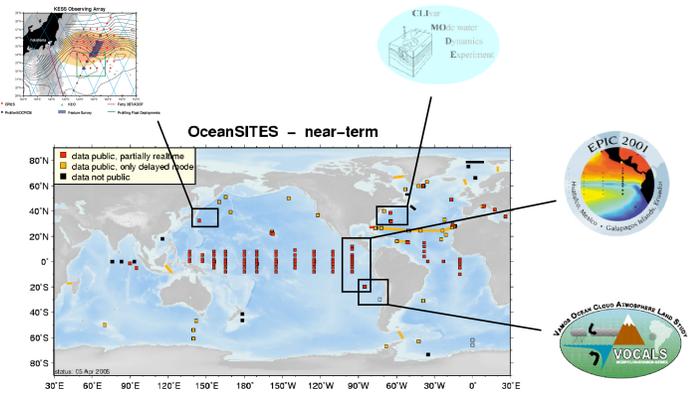


Fig. 1. The OCEAN Sustained Interdisciplinary Timeseries Environment observation System (OceanSITES) network of time series reference sites are indicated by squares. Red squares indicate surface meteorological sites. The boxed regions indicate U.S. CLIVAR process studies, including the Eastern Pacific Investigation of Climate (EPIC), VAMOS Ocean Cloud Atmosphere Land Study (VOCALS), Kuroshio Extension System Study (KESS), and the Climate Mode water, Dynamics Experiment (CLIMODE).



The global network of OceanSITES time series reference sites provides the longer-term monitoring context for process studies. In this poster, we use surface heat flux and cloud forcing "benchmark" data sets collected during the EPIC experiment to assess Numerical Weather Prediction (NWP) Reanalyses.

## Methodology

Net Surface Heat Flux:  $Q_0 = (1-a)SWR - \epsilon(\sigma T_s^4 - LWR) - Q_{lat} - Q_{sen}$

Latent Heat Loss (Qlat) and Sensible Heat Loss (Qsen) were computed with the Fairall et al. (2003) COARE v3.0 bulk algorithm using hourly data or telemetered daily averages when high resolution data were not available. These data included winds relative to surface currents, air temperature, relative humidity, and skin temperature. Bulk SST was extrapolated to the surface using the algorithm's warm layer and cool skin model. For more details on the latent heat flux calculation, see Cronin et al. (2006a)

Solar Cloud Forcing = downwelling shortwave radiation - clearsky shortwave radiation  
Longwave Cloud Forcing = downwelling longwave radiation - clearsky longwave radiation

Clearsky shortwave radiation was computed from Iqbal (1988) algorithm with decimal year/day, latitude, longitude, and estimates of integrated water vapor, and aerosol optical depths.

Clearsky longwave radiation was computed from the Fairall et al. (2006) algorithm with SST, specific humidity, air temperature, and latitude.

For more details on the surface cloud forcing calculation, see Cronin et al. (2006b).

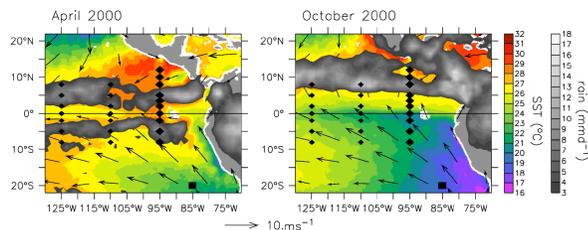


Fig. 2. The EPIC mooring array, shown in relation to the April 2000 and October 2000 mean TMI SST, TRMM rain rate, and QuikScat surface winds. Diamonds indicate TAO buoys. Large diamonds indicate EPIC-enhanced 95°W TAO buoys. The Woods Hole IMET buoy is indicated by a large square.

The objective of the Eastern Pacific Investigation of Climate Studies (EPIC) was to observe and understand the ocean-atmosphere processes responsible for the structure and evolution of the large-scale atmospheric heating gradients in the stratus deck / cold tongue / ITCZ complex, with the goal of improving short-term climate analysis and prediction systems for the Americas.

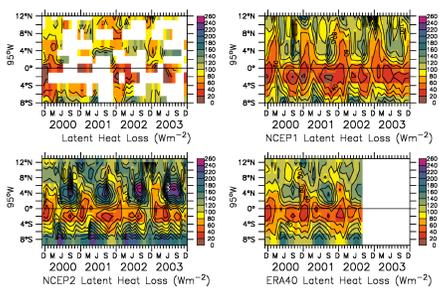


Fig. 3. Monthly averaged latent heat loss from (top left) 95°W TAO/EPIC buoy measurements, (top right) NCEP/DOE reanalysis, (bottom left) NCEP/DOE reanalysis, and (bottom right) ERA40 reanalysis

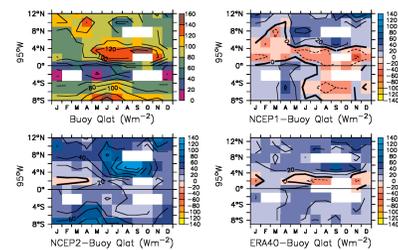


Fig. 4. (top left) Mean annual cycle of latent heat loss from buoy measurements, and mean annual cycle of difference between reanalysis latent heat loss and buoy field for (top right) NCEP/DOE, (bottom left) NCEP/DOE, and (bottom right) ERA40 reanalyses

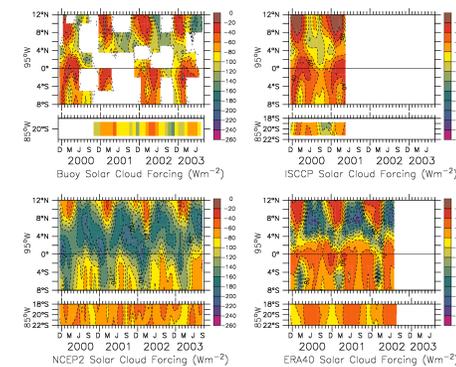


Fig. 5. Monthly averaged solar cloud forcing from (top left) buoy measurements, and (top right) NCEP/DOE, (bottom left) NCEP/DOE, and (bottom right) ERA40 reanalyses

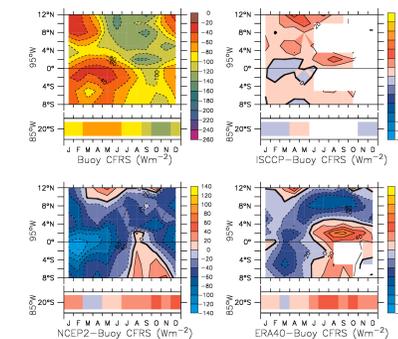


Fig. 6. (top left) Mean annual cycle of solar cloud forcing from buoy measurements, and mean annual cycle of difference between reanalysis solar cloud forcing and buoy field for (top right) NCEP/DOE, (bottom left) NCEP/DOE, and (bottom right) ERA40 reanalyses

- The TAO/EPIC 95°W buoys suffered extensive data loss due to vandalism.
- NCEP/DOE reanalysis tends to produce too much latent heat loss, particularly in the frontal region north of the cold tongue and in the region south of 5°S.

- ISCCP seasonal climatology compares relatively well with the buoys
- NCEP/DOE has too much cloudforcing on the equator during the warm season. During the cold season when stratus tends to form, there is too little cloud forcing in the NCEP/DOE and ERA40 reanalyses.

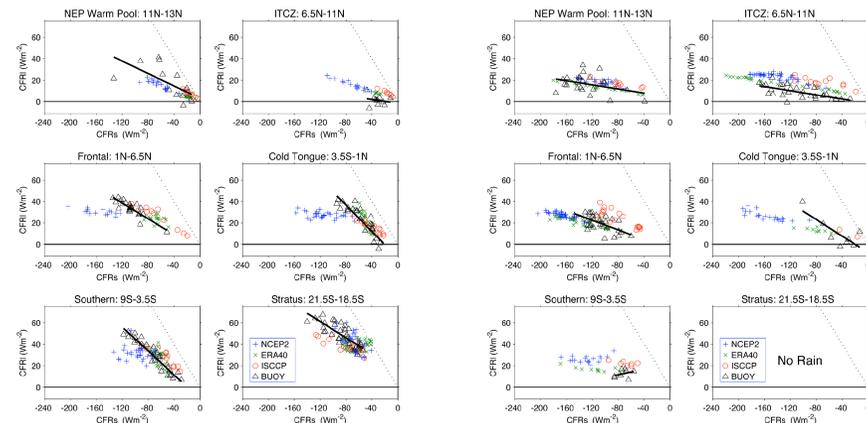


Fig. 7. Scatter plots of solar and longwave cloud forcing in different latitudinal bands for months with no significant rainfall. Significant rainfall was determined by TRMM rainfall fields.

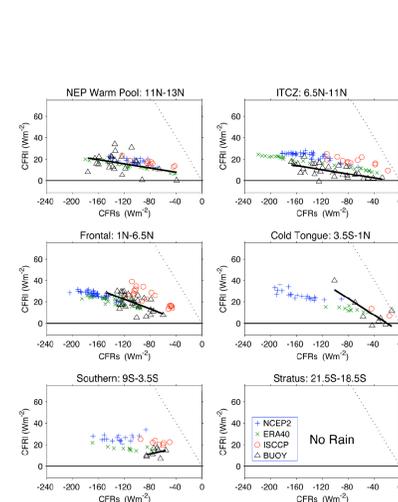


Fig. 8. Scatter plots of solar and longwave cloud forcing in different latitudinal bands for months with significant rainfall.

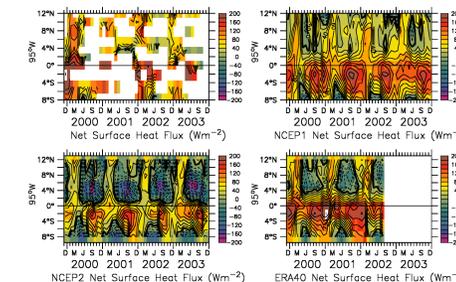


Fig. 9. Monthly averaged net surface heat flux into the ocean from (top left) 95°W TAO/EPIC buoy measurements, (top right) NCEP/DOE reanalysis, (bottom left) NCEP/DOE reanalysis, and (bottom right) ERA40 reanalysis

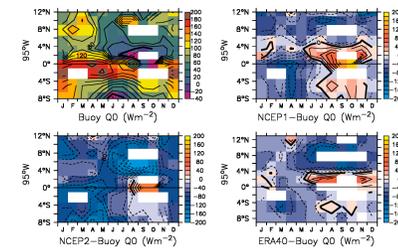


Fig. 10. (top left) Mean annual cycle of net surface heat flux from buoy measurements, and mean annual cycle of difference between reanalysis net surface heat flux and buoy field for (top right) NCEP/DOE, (bottom left) NCEP/DOE, and (bottom right) ERA40 reanalyses

- On average net surface heat flux warms the ocean everywhere within the eastern tropical Pacific, except in the frontal region north of the cold tongue during the cool season.
- All 3 reanalyses (NCEP/DOE, NCEP/DOE, and ERA40) do not have enough heat entering the ocean in the ITCZ regions.
- ERA40 and particularly NCEP/DOE reanalyses have net surface cooling in the northern ITCZ, rather than weak surface heat flux warming as observed.
- This is due to both improper cloud formation and cloud radiative balances, and to excessive surface latent heat loss.

## For more information

<http://usclivar.org/>  
<http://www.oceansites.org/>  
<http://www.pmel.noaa.gov/tao/epic/>

## References

Cronin, M. F., C. W. Fairall, and M. J. McPhaden, 2006a: An assessment of buoy-derived and numerical weather prediction surface heat fluxes in the tropical Pacific. *J. Geophys. Res.*, 11, C06038, doi:10.1029/2005JC003324.

Cronin, M. F., N. A. Bond, C. W. Fairall, and R. A. Weller, 2006b: Surface cloud forcing in the east Pacific stratus deck cold tongue / ITCZ complex. *J. Climate*, 19(3), 392-409.

Fairall, C. F., E. F. Bradley, J. E. Hare, A. A. Grachev, and J. B. Edson, 2003: Bulk parameterization of air-sea fluxes: Updates and verification for the COARE algorithm. *J. Climate*, 16, 571-0591.

Fairall, C. W., T. Uttal, D. Hazen, J. Hare, M. F. Cronin, N. A. Bond, and D. E. Veron, 2006: Cloud, radiation, and surface forcing in the equatorial eastern Pacific. *submitted to J. Climate*.

Iqbal, M., 1988: Spectral and total sun radiance under cloudless skies. In: *Physical climatology for solar and wind energy*, R. Guzzi and C. G. Justus, Eds. World Scientific, Teaneck, NJ, pp. 196-242.